Research Article

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Influence of diagenetic features on petrophysical properties of fine-grained rocks of Oligocene strata in the Lower Indus Basin, Pakistan

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Abstract: Nari Formation is considered as one of the most important oil and gas exploration targets. These finegrained tight sandstone reservoirs face enormous challenges due to their extremely low matrix porosity and permeability. Hence, in this regard, the study was carried out to collect the high-quality data on petrophysical properties along with mineralogy and microstructural characteristics and diagenesis. The experiments performed includes the petrographic study and scanning electron microscopy, and X-ray diffraction analyses. Besides, the measurement of petrophysical properties was carried out to assess the likely influence of the reservoir quality. The petrographic analysis shows predominantly fine- to medium-grained grey samples along with calcite, clay, lithic fragments and iron oxides. Further, the thin-section observations revealed that the quartz is a principal mineral component in all the analysed samples ranging from 52.2 to 92.9%. The bulk volume of clay minerals that range from 5.3 to 16.1% of. The porosity and permeability measured range from 5.08 to 18.56% (average 7.22%) and from 0.0152 to 377 mD (average 0.25 mD), respectively. The main diagenetic processes that affected the sandstones of Nari Formation are mechanical compaction, grain deformation, cementation and quartz dissolution and have played a significant role in influencing the quality of the reservoir rock. Overall, it appears that the primary petrophysical properties (porosity and permeability) were decreased due to the mechanical compaction, lithification, cementation, and framework grain dissolution. Based on the integrated mineralogical, microstructural analysis, and the laboratory-based petrophysical properties, the samples exhibited poor porosity, permeability, and moderate clay content, which indicate that the Nari Formation is a poor quality reservoir.

Keywords: tight sands, diagenesis, reservoir quality, petrophysical properties, Nari Formation, Lower Indus Basin Pakistan

1 Introduction

In today's world, the dominant source of energy being used by industrial sectors is mainly the hydrocarbons [1-4,41]. However, due to urbanization and industrialization, the demand for energy has rapidly increased [1,2,42]. The present-day demand of energy requirements cannot be met by the conventional resources. Therefore, oil and gas industries have made efforts to develop the unconventional hydrocarbon resources (tight gas) [1–5,43]. It is very challenging to develop such resources by conventional means (this is due to the tight formation exhibiting extremely low matrix porosity and permeability and possessing complex pore network) and it involves advanced technology, i.e. acidizing [3,44-46], hydraulic fracturing [2-10], and liquid nitrogen fracturing (recently developed technique) [9-12]. Fine-grained sedimentary formations are typically tight and are considered as nonporous or impermeable reservoirs with less than 10% matrix porosity and 0.1 milli darcy (mD) permeability [12,47]. Production from such hydrocarbon resources (tight gas) is very challenging due to low matrix porosity and permeability, presence of clay mineralogy, stress sensitivity, and rock heterogeneity [1–6,48–50]. Hence, it is essential to develop efficient exploitation and

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development strategies for effective production from tight resources. Effective production from any reservoir depends on the reservoir's quality [1-7,51]. The key factors that influence any reservoir's quality include (a) petrophysical rock properties (porosity and permeability), (b) depositional environment, and (c) diagenesis [24-30]. In addition, the petrophysical and microstructural characterizations are essential to achieve effective production, storage, and fluid flow from unconventional hydrocarbon reservoir rocks [1,2,5]. Petrophysical properties (e.g. permeability, porosity, water and oil saturations) of these rocks are mainly controlled by diagenetic changes [2,6-12]. Therefore, detailed assessment of petrophysical properties, i.e. permeability, porosity, and mineralogical and microstructural characteristics, is essential and an integral part of any field development plans. Further, the reservoir's fluid flow properties are controlled by pore morphology and depositional environment, diagenetic modifications, geothermal gradients, and pore pressures [2,8-11]. Main reservoir rock properties that are affected by diagenetic events are the porosity and permeability, which in turn are partly affected by the textural characteristics such as grain size, shape, grain packing and their arrangements [4,8,10,12-14]. Subsequently, the diagenetic processes affecting fine-grained sedimentary rocks include the grain fracturing, deformation, compaction, dissolution, authigenesis, biological actions, hydration, cementation, and recrystallization [15-27]. Mostly, the petrophysical data sets are unavailable for tight formations of the Lower Indus Basin of Pakistan; hence, the exploration and production companies are in an initial phase and are relying on the data sets that are available in public domain for the assessment of unconventional prospects. An integrated study was carried out to collect comprehensive high-quality data on petrophysical properties, mineralogy, microstructure, and diagenesis of samples obtained from Lower Indus Basin, Pakistan.

The objectives of present study are threefold: (i) collect data sets of high-quality petrophysical properties, (ii) integrate the petrophysical properties and microstructural data sets to assess the reservoir quality and (iii) improve understanding of diagenetic events and their impact on reservoir rocks quality. For this reason, we obtained exposed outcrop analogue samples from Nari Formation Lower Indus Basin, Pakistan. For accurate estimation of rock properties, it is essential to have a set of whole core samples recovered from wells while drilling the reservoir section. Obtaining a fresh core from reservoir section is the best option to be used for reservoir characterization studies [20,28]. However, it is

very difficult to access such kinds of samples, because these are often unavailable or their presence is very scarce or most of the companies secure them for their advanced phases of the project. In such circumstances, it would be far better to get outcrop analogues that could be used to acquire preliminary evidence about target reservoir. The present study obtained outcrop analogues for describing the petrographic and petrophysical characterization and further described the diagenetic events that impact the quality of reservoir rocks and to collect the high-quality data on petrophysical properties of these sandstones.

1.1 Geological context of the study area

The Indus Basin is located between the boundary of the Indian, Asian and Arabian plates [29]. The Indus Basin is

Table 1: Generalized sequence stratigraphy along with main lithology encountered in the Indus Basin Pakistan [52]

A	ЗΕ	FORMATION (Member)	LITHOLOGY									
PLIOCENE MIOCENE		SIWALIKS										
OLIG		NARI										
Е		KIRTHAR	111111111111111111111111111111111111111									
EOCENE		GHAZIJ SULAKI										
PAL		RANIKOT	Summer									
		PAB MUGHAL KOT										
Sn	LOWER UPPER	PER	PER	PER	PER	PER	PER	PER	PER	PER	PARH	
CRETACEOUS		Upper GORU										
CRET		Lower GORU	70									
	LOV	SEMBAR										
	UP.											
JURASSIC	MIDDLE	CHILTAN										

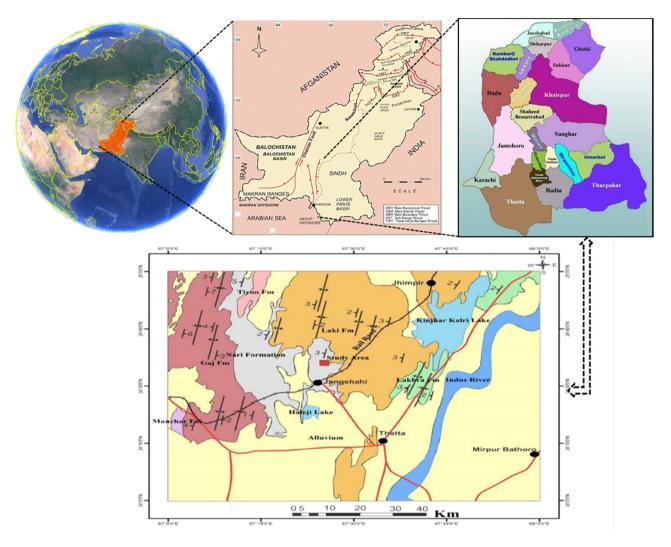


Figure 1: Geological map of Pakistan showing the study area with basic geological information and location of sampling sites.

divided into two parts, Upper Indus Basin and the Lower Indus Basin [30,31]. The general stratigraphy of Lower Indus Basin including various different formations is presented in Table 1. The study area lies in the Karachi trough, part of Lower Indus Basin. The rocks are ranging from Eocene to subrecent. The lower contact of mostly exposed Nari Formation in Sindh is mostly conformable and gradational with Kirthar Formation while in our study area it is unconformable with Laki Formation and upper contact is unconformable with recent deposits. The formation consists of sandstone, shale, and subordinate limestone. The study area lies in the upper part of Nari Formation where limestone is absent; generally fossiliferous limestone is present around Karachi City [31,32]. It is mainly composed of sandstone with minor amount of shale [32] (Figure 1).

2 Sampling and methodology

This study addresses the petrographic and petrophysical characterizations of potential oil and gas exploitation target in the region. The overall methodology to analyse the petrographic and petrophysical characteristics of Nari Formation is depicted in Figure 2, and the number of samples analysed and the measurements are given in Table 2, considering the following analysis steps.

- Selecting the study area based on the potential reservoir rock
- Identifying the surface outcrop as an analogue to characterize the reservoir rock
- Collecting and preparing samples for laboratory analysis
- Analysing of petrographic thin section and mineral contents

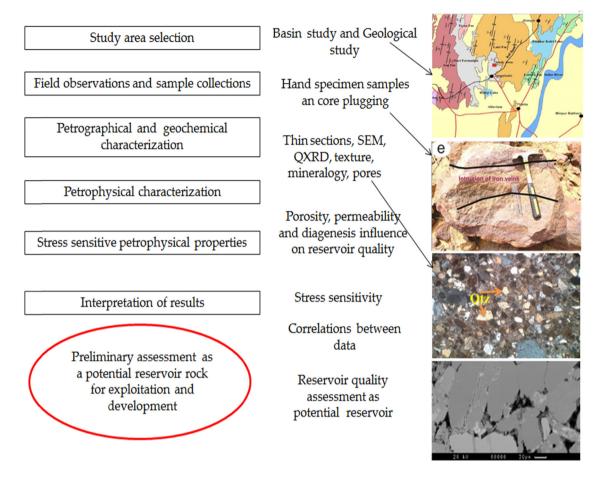


Figure 2: The schematic representation of the overall workflow followed in this study.

- Measuring petrophysical properties
- Interpreting data and correlating petrographic and petrophysical data for reservoir quality assessment.

2.1 Petrography

For petrographic thin-section observations, Leica optical microscope was used to study the rock lithology and microstructure of the Nari sandstone samples. The experiments were carried out at the Geological Survey of Pakistan (GSP), Karachi laboratory. In addition, cathodoluminescence can identify minerals and fossils, along with their microscopic features. The Leica optical microscope was used to analyse the essential microscopic features of thin section under polarized microscopes.

2.2 XRD analysis

X-ray powder diffraction (XRD) is a fast analytical method basically adapted to identify the different phases of crystalline materials. Prior to analysis, the material was powdered, finely crushed, and made very homogeneous so the average bulk composition of material is obtained. These experiments were conducted at the University of Sindh Geology Department. QXRD analysis was performed on 1 cm³ size, and the sample covering the overall sandstone mineralogy was also taken.

2.3 Scanning electron microscopy (SEM) analysis

The smooth and fresh rock chips of samples were prepared and then mounted on the sample stub of SEM. The samples were mounted with the help of the conductive double-sided carbon solution tape. The sample stub was then mounted in the sample chamber JEOL JSM-6590LV of SEM, which is also equipped with an extra and important accessory component, Bruker of energy dispersive X-ray spectrometer (EDS). The IMAGINING INTERFACE software controlled the system in JEOL JSM-6590LV in

Table 2: Summary of the experiments carried out and samples collected from Nari Formation in the Lower Indus Basin, Pakistan

S. No.	Sample ID	XRD	Petrography	SEM- EDS	Porosity and permeability
1	18-JT-1a	$\sqrt{}$	$\sqrt{}$	\checkmark	\checkmark
2	18-JT-2		\checkmark	\checkmark	\checkmark
3	18-JT-3		×	\checkmark	\checkmark
4	18-JT-5	$\sqrt{}$	\checkmark	\checkmark	\checkmark
5	18-JT-6a		\checkmark	\checkmark	\checkmark
6	18-JT-6b	$\sqrt{}$	×	\checkmark	×
7	18-JT-4		\checkmark	\checkmark	\checkmark
8	18-JT-7		\checkmark	\checkmark	\checkmark
9	18-JT-8	$\sqrt{}$	\checkmark	\checkmark	\checkmark
10	18-JT-9	×	×	×	\checkmark
11	18-JT-10	$\sqrt{}$	\checkmark	\checkmark	×
12	18-JT-10a	×	×	×	\checkmark

SEM whereas in the EDS, QUANTAX. The desired images were taken and controlled from computer to study the various topographic features and compositions of the samples.

2.4 Measurement of density, porosity, and permeability

Cylindrical samples were prepared for permeability and porosity measurement. The samples prepared in varying lengths and diameters. All samples were cleaned to remove the unnecessary particles which may affect results. All the samples were then dried into a humidity-controlled oven at 60°C prior to measurements.

2.4.1 Bulk density measurement

The bulk density of rock samples was obtained by determining the core plugs (cylindrical plugs), dimension, and the bulk volume (V_b) ; and the dry weight of these samples (W_d) was obtained using a precision calliper (0.1 mm precision) and an electronic balance (0.1 mg precision).

2.4.2 Porosity and permeability tests

Permeability and porosity measurements were made for the samples obtained and the experimental procedures are described in ref. [7].

3 Results

3.1 Petrographic analysis

Thin-section observation shows that Nari Formation samples are moderate to poorly sorted. The mineralogical compositions are shown in Table 3, and the diagenetic features are shown in Figure 3. The occurrence of quartz considerably varies in proportion from 62 to 91.20%. The observed accompanying mineral assemblage includes the calcareous matrix, ferric oxides, calcite cement, clays, and lithic grains along with feldspars and mica. The proportions of all the associated minerals except quartz are shown in Figure 4. Ferric oxide, calcareous mud, and calcite observed from thin-section analysis dominated the cementing materials. These cements very densely

Table 3: Mineralogical composition from petrographic thin sections point counting analysis

S. No	Sample ID	Mineral components (%)						
		Quartz	K-feldspars	Plagioclase	Mica	Calcite	Lithic fragments	Clay mineral
1	18-JT-1a	77	12	3.0	2.0	4.0	4.0	2.0
2	18-JT-2	83	5.0	8.0	2.0	2.30	2.0	3.0
3	18-JT-3	76	7.0	5.0	2.0	2.00	3.0	3.0
4	18-JT-5	89	2.0	2.0	3.0	1.00	1.0	2.0
5	18-JT-6a	69	8.0	5.0	4.0	5.00	6.0	3.0
6	18-JT-6b	90	4.0	3.0	2.0	3.00	1.0	_
7	18-JT-4	75	9.0	3.0	3.0	5.00	3.0	2.0
8	18-JT-7	73	7.0	5.0	2.0	3.00	4.0	9.0
9	18-JT-8	75	3.0	3.0	2.0	3.0	2.0	15.0
10	18-JT-10	91	6.0	_	2.0	_	2.0	12.0

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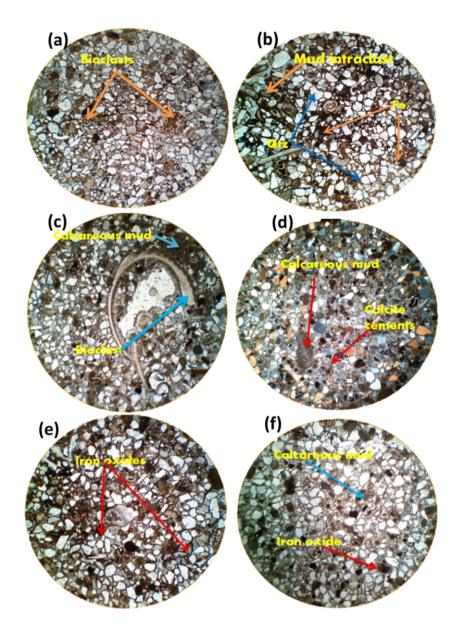


Figure 3: Thin-section photomicrograph of Nari Formation sandstones showing diagenetic features: (a) Bioclasts. (b) Mud interclasts along with iron oxides among quartz grains. (c) Calcareous mud and bioclasts are shown. (d) Calcite cements and calcareous muds are visible from the image. (d) Iron oxides coating quartz grains are visible from the image. (f) Calcerous muds and iron oxides are shown.

packed these rocks, thereby resulting in poor pore spaces and subsequently leading to tight sandstones which exhibit poor porosity and permeability. From Figure 4, it was observed that the clay is the dominating mineral in samples 18-JT-8 and 18-NR-2. Many observations showed that the monocrystalline quartz was in principal quantity than the polycrystalline. At most places, the quartz grains are fractured and it may be due to the external stress (Figure 5b). The tectonic stress could be the major cause of grain breakage and fracturing (Figure 5d). The dark coloured patches within these samples display those

iron oxides within these Nari sandstone samples along with calcareous mud. Both alkali feldspar (Figure 5c) and plagioclase (Figure 5c) were observed in the samples.

3.2 Microstructural observations

Samples selected for the SEM-EDS analysis were named after their location from where they were collected, such as 18-NR-1 and 18-NR-2 were collected from station 1, and those which were collected from station 2 were labelled

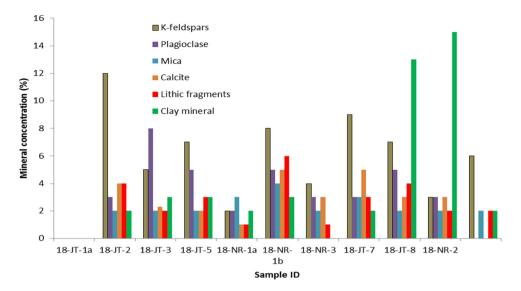


Figure 4: Mineral concentration of samples of Nari Formation other than quartz, as quartz is a highly dominated mineral within this formation.

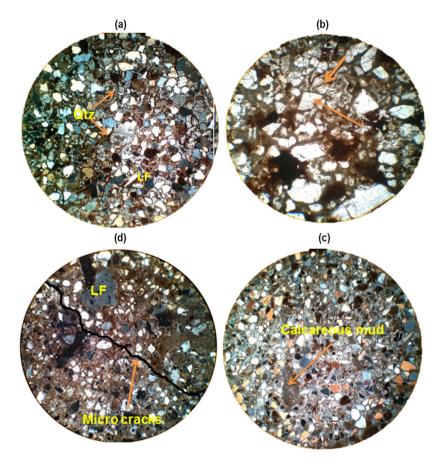


Figure 5: Photomicrographs of sandstone samples analysed from Nari Formation showing: (a) moderately sorted, subangular to subrounded grains of quartz. (b) Monocrystalline and polycrystalline characteristics governing within these samples as quartz crystal. (c) Lithic framework grains along with observed fracture. (d) Some observed alkali feldspar grains.

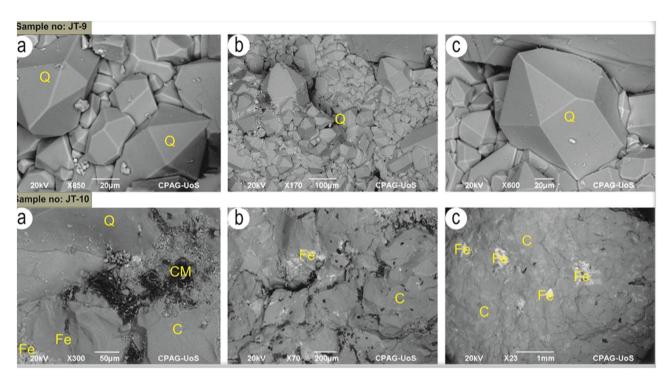


Figure 6: Mineral composition under SEM observation is quartz and calcite (in Figure a, b, and c sample JT-9). Quartz mineral indicates its tone as dark grey while the calcite mineral indicates its tone as light grey. The sample JT-10 (a-c) displays quartz as its dominant mineral along with iron oxides, calcite, and clay minerals. Clay minerals as kaolinite are present in JT-10 (a). The symbols represent Q = Q quartz, Q = Q calcite, Q = Q minerals.

as JT-1, JT-2, 3, 4, 5, 8, 9, 10, 12, and 13. From SEM analysis, the various minerals identified are represented as Qtz = quartz, C = calcite, Fe = iron, and CM = clay minerals.

All samples from stations 1 and 2 were analysed in order to identify their general elemental structure. In JT-1 sample, the general concentration of elements is analysed under EDS in the form of peaks. The highest peaks of this sample are of silicon, oxygen, calcium, and iron, as shown in Figures 6 and 7. The presence of dominant minerals such as quartz, calcite, and iron sand was confirmed under SEM analysis of sample 18-JT-1. The point analysis of this sample is based on the white colour. Figure 7 indicates the highest peaks of the silicon, iron, and calcium in which the tone of the iron was confirmed under SEM analysis of sample 18-JT-1. Sample 18-JT-2 general elemental concentration indicates the highest peak of silicon, and all other peaks are averaged as shown in Figure 7. The presence of quartz was confirmed under SEM analysis of sample 18-JT-2. Further elemental concentration of sample 18-JT-3 indicates the highest peaks of silicon, oxygen, and calcium in Figure 7. The presence of quartz and calcite minerals were confirmed under SEM analysis of sample 18-JT-3. The point analysis (a) of this sample is based on the dark grey shown in

Figure 7, which indicates the highest peak of silicon7. The tone of the quartz was confirmed under SEM analysis of sample 18-JT-3. The point analysis of this sample is based on the light grey indicating the highest peaks of the oxygen, silicon, and calcium (Figure 7). The tone of the calcite was confirmed under SEM analysis of sample JT-3. Further, it was found that in the samples 18-JT-4 to J18-T-8, their elemental concentration ranged from their highest peaks of silicon, oxygen, and calcium, and the presence of quartz and calcite minerals were confirmed under SEM analysis.

The general elemental concentration of 18-JT-9 sample indicates the highest peak is of silicon and then oxygen as the lowest peak as shown in Figure 7. The presence of the dominant mineral quartz was confirmed under SEM analysis of sample JT-9, which consists of beautifully developed crystals of quartz (Figure 6). The general elemental concentration of 18-JT indicated the highest peak of silicon and oxygen. The presence of dominant mineral quartz was confirmed under SEM analysis of sample JT-10. The point analysis (Figure 7) of this sample is based on the dark black seen in the image which indicates the highest peaks of silicon, oxygen, potassium, and calcium. The tone of the clay minerals was confirmed under SEM analysis of sample JT-10.

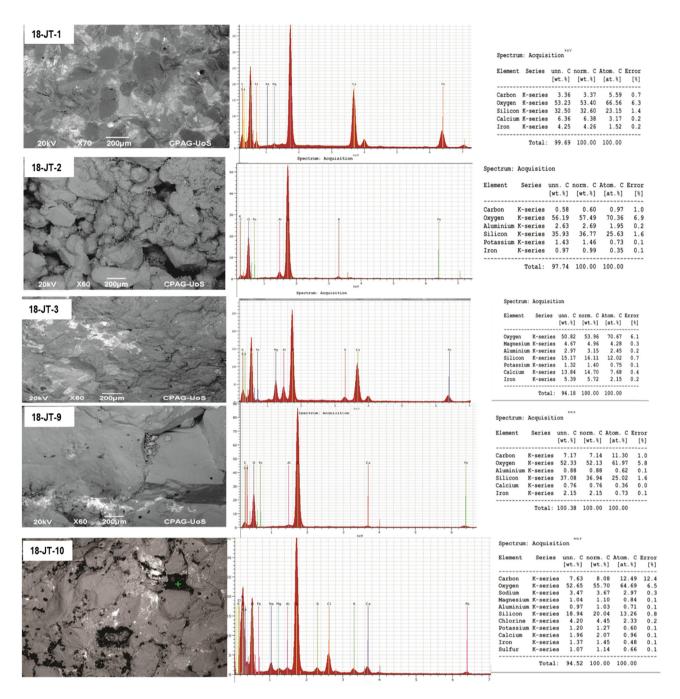


Figure 7: The scanning electron microscopy images of Nari sandstone samples along with energy dispersive X-ray (EDX) spectrum revealing the different elemental composition.

3.3 Mineral composition

Results from XRD analysis are presented in Table 3; this shows the compositional analysis for all analysed sandstone samples. It was observed that the main components are quartz and feldspar, and the iron oxide, cement, and clay of variable mineralogy. Other minerals observed include the kaolinite, dolomite, calcite, gypsum, mica, pyrite, chlorite,

and lithic fragments. The cementing materials observed are mainly comprised of carbonate (9.4%), ironic (10.6%), and micrite muds (12.8%); in the samples collected from Nari station 1, the major cementing materials are clay minerals (14.2%), ironic (4.7%), and plagioclase (2.2%). From the XRD analysis (Figure 8), it was found that the quartz content in all samples is very high, i.e., it was 91.2% and above 91.6%, which is typically quartz sandstone. Most of the samples are

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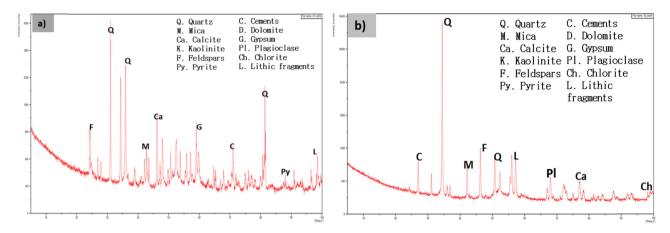


Figure 8: The illustration shows the X-ray diffraction patterns of two selected sandstone samples (a and b) from Nari Formation.

Table 4: The XRD results signifying the mineral composition and concentration (wt%)

Sample ID	Mineral components (%)							
	Quartz (SiO ₂)	K-feldspars	Plagioclase	Mica	Calcite	Lithic fragments	Clay mineral	
18-JT-1a	77	8.20	3	2	4	4	2	
18-JT-2	83	5.30	2	_	2	5	3	
18-JT-3	76	9.10	3	2	_	9	2	
18-JT-4	69	12.3	4	2	3	8	2	
18-JT-5	75	9.12	3	2	4	4	2	
18-JT-7	63	5.27	3	4	3	6	12	
18-JT-8	65	3.12	2	3	3	4	16	
18-JT-9a	71	10.02	5	2	4	6	2	

lithic quartz sandstone, possessing clay minerals in variable quantities as shown in Table 3. The feldspar content is very small and can be neglected. The average total clay content (24.35%) of Station 1 Formation is larger than the average total clay content (15%) in Station 2 Formation in the region (Table 4).

3.4 Petrophysical property results

Permeability and porosity data of Nari sandstones along with their slippage factor values are presented in Table 5. The results obtained were cross plotted and are shown in Figure 9, where a reasonably virtuous correlation between permeability and porosity can be perceived on the semilogarithmic plot. It is interesting to note that, the permeability value >0.1 mD, shows the linear correlation between permeability and porosity. However, the permeability and porosity values <0.01 mD are quite scattered on the semilog plot, which means that the effect of other parameters such as pore-filling clays and

grain-coating clays along with cementing materials significantly affect the permeability. The cross-plotted permeability and porosity demonstrate that the trend line of samples with no clay minerals shows linear trend line; however, the sandstones possessing clay minerals, calcareous muds, and cements resulted in low and scattered permeability and porosity, which in turn makes this reservoir very poor quality. The permeability varies by several factors due to the varying mineral proportions. The main factors that control the permeability of these reservoir rocks are clay mineralogy, cementing and mud influx materials. The permeability reductions experienced imply that the

Table 5: Basic petrophysical properties of Oligocene Nari sandstone in the Lower Indus Basin Sindh, Pakistan

	Petrophysical parameters	Min	Max	Mean	SD
1	Rock density (g/cc)	2.27	2.69	2.390	2.20
2	Porosity (%)	5.80	18.5	7.20	3.30
3	Permeability (mD)	0.0152	377	0.052	32.2
4	Gas slippage factor (1/psi)	5.0	125	43.50	12.5

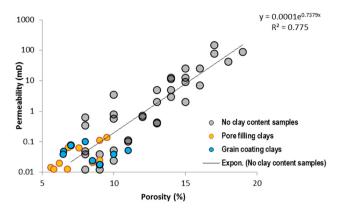


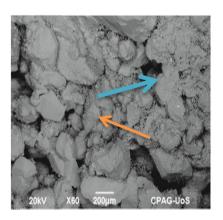
Figure 9: The cross plot of porosity and permeability of samples from Nari sandstone formation.

Nari Formation could have a significant impact on the single-phase flow, especially when production from these reservoirs begins. Further the samples which were cemented due to diagenetic modifications have shown large reduction in permeability values. It was observed that some low-permeability sandstones exhibited overlapping permeability-porosity characteristics (Figure 10). Almost identical behaviour of permeability data could be noticed by the action of grain-coating clays and porefilling clays (Figure 10). The overlapping was usually noticed in the data where permeability of these samples was very low, that is, less than <0.1 mD particularly for rocks with pore-filling clays and grain-coating clays. The distribution characteristic of clay mineralogy within sandstone formation is described in Figure 10. The lower permeability of Nari Formation sands is the result of the existence of clays, which is obvious from SEM and XRD data. Pore-fillings clays demonstrate that the permeability in these sands is less compared to the graincoating clays; however, this needs further refinement for accurate prediction of permeability and porosity of Nari Formation reservoirs. Laboratory measurements of porosity and permeability of numerous sandstone and mudstone samples have demonstrated that it is mainly controlled by the clay mineralogy. From the permeability and clay content cross plot, a scatter of data is observed (Figure 11). The scatter in permeability and porosity could be the result of different grain sizes and occurrence of diverse mineral contents and is well elaborated in ref. [33].

4 Discussion

4.1 Diagenetic controls on quality of reservoir

The main diagenetic processes that affected the Nari Formation sandstones are the mechanical compaction, deformation, cementation, and dissolution, which play a significant role in the reservoir quality of the sandstone. It has been observed that the porosity and permeability reduction as a result of calcite cements are larger than those due to the coating cement (i.e. clay minerals). It has been reported that the calcite cement observed in the Nari sandstone formation usually exists in other sandstone formations around the globe e.g. ref. [34]. Subsequently, it is perceived that the fluid flow may exhibit difficulty in moving through pores with a higher percentage (>20%) of calcite cements [35-37]. On the other hand, if calcite cements are found to be in smaller amounts, that is, less than 20%, it may result in slight increase in secondary porosity [36,38]. The likely sources of calcite cements may consist of the carbonate rock fragments dissolution, dissolution of K-feldspar, the diagenesis of interclasts



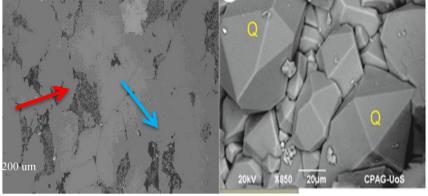


Figure 10: The clay mineralogy distribution and its impact on reservoir rock properties. From left to right this illustration displays the grain-coating clays, pore-filling clays, and no clay as observed from the scanning electron microscopy images.

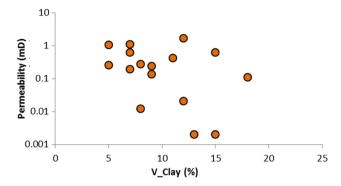


Figure 11: The relationship between permeability and clay content.

and calcareous muds, alteration of clay minerals such as kaolinite to illite, and the transformation of smectite to illite [4,34,38,39]. The Nari Formation sandstones mostly contain interclast muds, which leads to the likely chances of alteration of authingenic minerals. The petrographic analysis displays that calcite cements are dispersed in most of the places in the Nari Formation sandstone samples. Consequently, the key cause of calcite cements in these samples is believed to be the results of exterior mudstones and interclasts. The diagenesis that affected Nari Formation sandstones is mainly classified as early and late diagenesis (Table 6).

4.2 Paragenetic sequence

The textural connections among the diagenetic minerals combined with the petrographic studies were conducted to establish an overall paragenetic sequence of the Nari Formation. The main diagenetic minerals identified within these sample paragenesis are predominantly the cementation by carbonate mud influx, clay minerals, and quartz cements. Diagenetic modification showed discrepancy between different samples. The diagenetic events experienced by sandstone samples of Nari Formation is summarized in Table 6.

4.2.1 Early diagenesis

Depositional facies and detrital composition control the earlier processes of the diagenesis. They include mechanical compaction, earlier cementation dissolution, grain coating of smectite clay, and kaolinite precipitation. Porosity and permeability are controlled as inverse relation with mechanical compaction. The higher the compaction, the smaller the porosity and permeability in the sandstone facies. Studied samples indicate the mechanical compaction is mostly inferred by the closer grain packing of quartz grains (Figure 3a-f). Other indicators of mechanical compaction observed were the ruptured nature of earlier bioclasts, rock fragments, and mud intraclasts (Figure 3a-c). Cementation largely comprises iron oxide, early calcareous mud, and later stage calcite in the studied samples. The calcareous mud indicates the carbonate influx in the basin (Figure 3d and e). The cementation of iron oxide and calcareous mud also reduces the porosity and permeability in the studied samples significantly.

4.2.2 Late diagenesis

Detrital fragments such as feldspar undergo dissolution with interaction of meteoritic water. The meteoritic water come into contact with sediments when they are uplifted to shallower depths mostly of less than 2 km. The resulting dissolution alters the silicate such as feldspar and produce kaolinite. The degree of feldspar kaolinitization varies broadly. As such the fluvial sandstone alters more extensively than the shallow marine sandstone. Dissolution enhances the porosity and permeability. The presence of minor amount of kaolinite in the studied samples indicates lesser degree of kaolinitization (Figure 6 JT-10a) which shows smaller dissolution; hence, the trivial porosity and

Table 6: Illustrates the diagenetic events ex	perienced by Nari Formation sandstone reservoirs
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Diagenetic sequence	Eodiagenesis	Mesodiagenesis	
Temperature / °C	70	140	
Mechanical compaction			
Grain deformation			
Feldspars dissolution			
Quartz dissolution			
Carbonate cementation			
Smectite clay			
Kaolinitization			

permeability enhancement. Late stage calcite cementation is also evidenced in the samples.

which shows that minor dissolution process has occurred. Hence, the secondary porosity is not significantly produced.

4.3 Burial diagenetic modifications on reservoir quality

For the successful exploitation and development of any reservoir, it is essential to have a sound knowledge of their mineral compositions, microstructural characteristics, associated depositional environment and burial diagenetic history, overburden stress, and reservoir rock heterogeneity [13,14,40–51]. The main diagenetic processes that affected the Nari Formation sandstones are the mechanical compaction, deformation, cementation, and dissolution, which play a significant role in the reservoir quality of the sandstone. The following events played a significant role in controlling the porosity and permeability of the Nari sandstone.

4.3.1 Mechanical compaction

Mechanical compaction reduced the porosity by rearrangement, grain packing, and displacement of detrital sediments. Small amount of earlier diagenetic cement in the samples rendered compaction to be a dominant factor to reduce the pore spaces between the sediments. This resulted in the overall porosity and permeability reduction.

4.3.2 Cementation

In the studied samples, cementation mainly comprises iron oxide (Figure 3e), calcite, and early calcareous mud. This influx of calcareous mud indicates the deltaic to shallow marine depositional environment. The calcareous influx also brought some bioclasts (Figure 3a and c). The cementation of iron oxide and calcareous mud is of pore-filling type and significantly reduced the porosity and permeability in the studied samples in the later stage of diagenesis.

4.3.3 Dissolution

In the later stage of diagenesis, epidiagenesis is the main phase for the formation of secondary dissolution pores. Kaolinitization is less evident in the studied samples,

5 Conclusions

Integrated analyses of mineralogy, microstructure, petrography, and petrophysical properties on Nari sands were performed to evaluate the impact on the reservoir quality of the Lower Indus Basin, Pakistan. The following conclusions were drawn from the data analysis:

- (i) Analysed samples reveal that these are mainly composed of medium- to fine-grained, angular to subrounded quartz, associated with feldspars as well as some lithic fragments along with the broken fragments of fossils, cementing materials of iron oxides and calcites. The XRD result also confirms that quartz is the principal mineral component, forming 52.2–92.9% in all the analysed samples.
- (ii) XRD results demonstrate that the bulk volume of the samples possesses clay minerals ranging from 5.3 to 16.1%. The most common clay mineral in all the samples was chlorite forming 4.1–9.4% of the bulk volume, and the illite and kaolinite clays were in very small quantity and considered as insignificant.
- (iii) The Klinkenberg-corrected gas permeability obtained at maximum overburden stress of 2,500 psi confining pressure was ranging from 0.542 to 0.0112 mD. The presence of calcite and iron oxides along with clay minerals leads to significantly lower porosities and permeabilities resulting in poor quality reservoir.
- (iv) The main diagenetic processes that affected the Nari Formation sandstones are the mechanical compaction, deformation, cementation, and dissolution. The permeability and porosity of Nari sandstones are smaller in values due to the mechanical compaction, authigenic cements pore filling of minerals and cements such as mud interclass bioclats, clay minerals, and iron oxides which turned this into tight reservoir rock.

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