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Lithopetrographic and geochemical features of the Saalian tills in the Szczerów outcrop (Poland) in various deformation settings

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Abstract: This paper presents the results of new studies of Saalian tills, from the Ławki and Rogowiec formations, filling the Kleszczów Graben. The study area is located in the Szczerów outcrop, Bełchatów Lignite Opencast Mine, central Poland. Laboratory studies included macrofossil analysis of the deposits, as well as petrographic and geochemical (neodymium isotope ratio) measurements. The studies were carried out in 2014–2016 and resulted in both establishing the sedimentary Saalian complex and constructing geological cross-sections and a synthetic lithostratigraphic profile. Development of sediments in this part of the Kleszczów Graben in the Pleistocene was largely influenced by tectonic factors (Chabielice fault, Dębina Salt Dome) and glaciectonic processes (Wartanian Glaciation). The Saalian tills (T4 – Ławki and T7 – Rogowiec lithotypes) are between the Holsteinian sand with macroremnants of trees and the Eemian gyttja and peat. Petrographic coefficients for lithotype T4 (Ławki Formation) are 1.33–1.06–1.01 and 0.12 and for lithotype T7 (Rogowiec Formation) are

1.29–1.23–1.04 and 0.12. Investigations of the Nd isotopic compositions of the studied samples reinforce our interpretations of till deformation and tentative lithostratigraphic correlations. Neodymium isotope ratios “P” (-14.4 ± 0.7) and “D” (-12.4 ± 0.3) ε_{Nd} values (2σ) correspond to mean signatures of Saalian glacigenic moraine sediment. This proves that ε_{Nd} values less than -12.4 reflect the southern Fennoscandian Ice Sheet sediment provenance, but ε_{Nd} values greater than -12.4 indicate the western European origin.

Keywords: Pleistocene, Saalian, stratigraphy, glacial tills, petrography, Neodymium isotope, Szczerów outcrop, central Poland

1 Introduction

As the lignite in the Bełchatów and Szczerów outcrops is extracted using the opencast method (Figure 1), we are capable of tracing till horizons and their positions in the vertical section of Pleistocene deposits.

The investigations carried out during the last years have shown that the lithostratigraphic and chronostratigraphic subdivisions of the Bełchatów outcrop [1–4], which are also used for the Szczerów outcrop [5], do not correspond fully to the new subdivision established for the Polish Lowlands [6–8].

The lithology of Saalian deposits in the study area is relatively diverse. As the mining operations were progressing to the east, more and more intense effects of tectonic and glaciectonic deformation of Pleistocene deposits were observed in the Szczerów outcrop. This is because of both the occurrence of the Chabielice fault and the influence of the Dębina Salt Dome (DSD) [5,9,10]. About 10–30 km to the south and south-east of the study area, there is the limit line of the Wartanian Glaciation [11–13] or the Wartanian Stadial [14,15]. The study area and adjacent areas (Szczerów Basin and the southern part of the Bełchatów Plateau) are considered as examples of classical development of ice-lobe relief of the Widawka lobe [13,14,16,17]. In the regions of Biała, Sulmierzyce,

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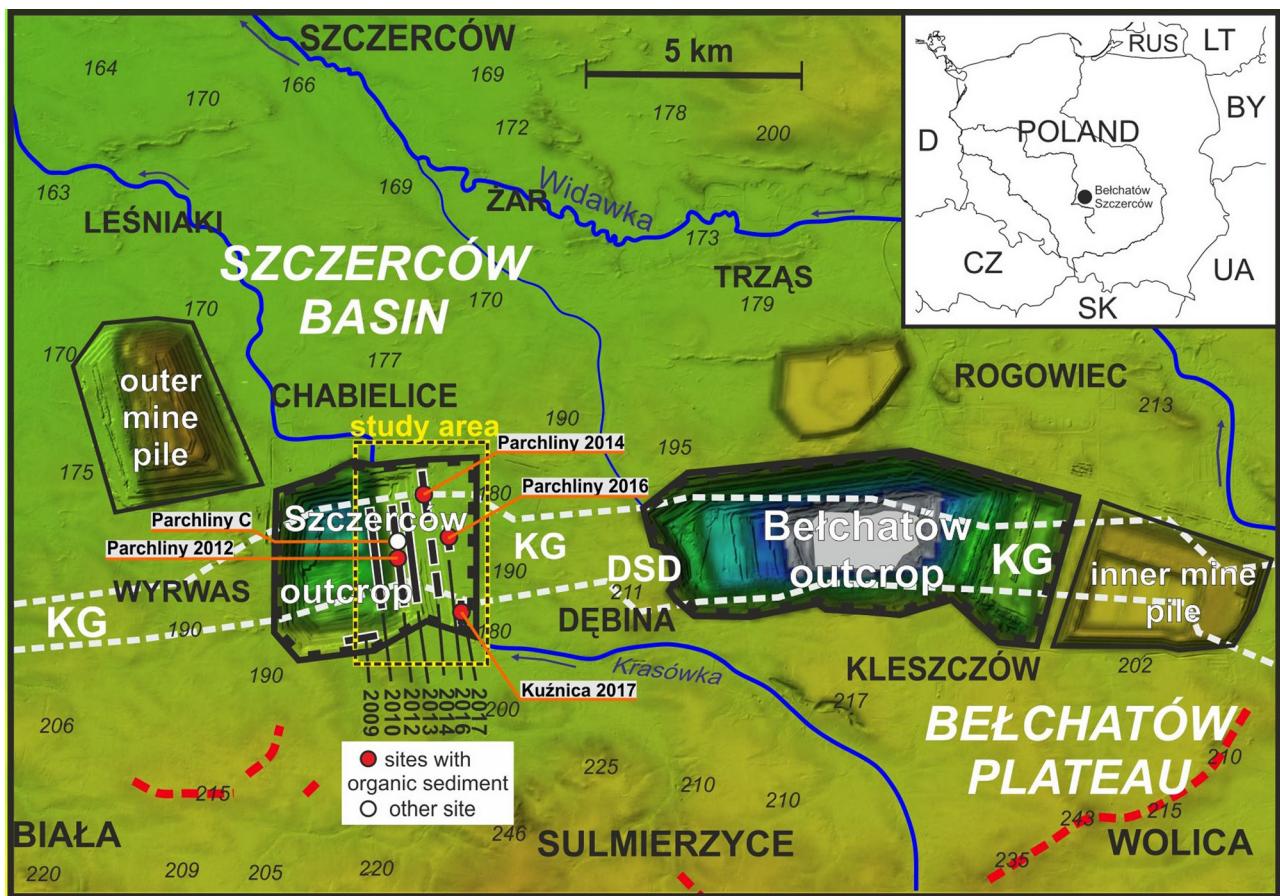


Figure 1: Study area against the background of the Belchatów Lignite Opencast Mine facilities and the surrounding relief. Elevations are in italics. Site with organic sediment: 2012 – “Parchliny 2012” (Early to Middle Eemian) and Parchliny C (Early Saalian); 2014 – “Parchliny 2014” (Eemian–Early to Middle Weichselian); 2016 – “Parchliny 2016” (Late Saalian to Early Eemian); KG – Kleszczów Graben (marked by white dashed line); DSD – Dębina Salt Dome.

and Wolica, there is a zone of frontal moraines, whereas toward the north, a zone of kettle lakes and holes extends (Figure 1).

2 Geological setting

Both (Belchatów and Szczerców) outcrops are located within the Kleszczów Graben, which was active in the Neogene and Pleistocene [18]. Although a wide range of Pleistocene sections should be preserved in this area, they are not present. On one hand, the ice sheet advance and the stress that the ice exerts on the bedrock led to subsidence of the graben's floor, which favored accumulation of thick ice-dammed lake or glaciofluvial series [19] in two to three cycles [20]. On the other hand, mining operations in the Szczerców outcrop indicate a local significant reduction of the Pleistocene section. This is

probably related to the location of the Chabielice fault and the DSD in this region [21].

In the last years, it was possible to make a comprehensive study of the following units: the Piaski Formation from the Middle and Late Weichselian [22], Aleksandrów Formation from the Late Saalian, Eemian Interglacial and Early Weichselian [23], and the Saalian complex: the Rogowiec Formation (T7, Wartanian) and Ławki Formation (T4; Odranian) [5,10,24,25] (see Table 1).

Tracing the horizons of mainly Saalian tills, from the Krasówka valley in the center of the outcrop toward the salt diapir (SDS) to the east, and against the Chabielice fault in the western sector of the Kleszczów Graben, was of foremost importance. Next, the role of tectonics, glaciotectonics, and halotectonics in the evolution of the graben's infill were studied in relation to the critical views and interpretation of the lithostratigraphic subdivision applied for the central part of the graben – the Belchatów outcrop.

Table 1: General stratigraphy of the Pleistocene deposits filling the Kleszczów Graben according to Krzyszkowski et al. [5,10,22,23,25]

Tectonics	Lithostratigraphy				Chronostratigraphy								
	WIDAWKA FORMATION		SZEROKIE FORMATION		Szczerców outcrop	Bełchatów outcrop	HOLOCENE						
UPPER STRUCTURAL UNIT (UNDISTURBED)	PIASKI FORMATION		3	a		14 500 BP	WEICHSELIAN	UPPER					
				b	43 500 BP	21 000 BP		MIDDLE					
				c	24 080 BP 33 090 BP	27 000 BP							
				d		33 000 BP							
				e	45 000 BP > 47 000 BP	> 43 000 BP							
	ALEKSANDRÓW FORMATION				Eemian Interglacial		EEMIAN	Early					
					TILL 7								
	ROGOWIEC FORMATION (glacial)				TILL 7		SAALIAN	WARTANIAN 2					
	A	?			interstadial?								
	B	TILL 6			WARTANIAN 1								
LOWER STRUCTURAL UNIT (DISTURBED)	CZYŻÓW FORMATION			C	?			PILICIAN					
				A	TILL 5			?					
				B	PILICA INTERSTADIAL			ODRANIAN					
					TILL 4			?					
					TILL 4								
					?								
					TILL 3								
				A	Podlesie Interstadial		CZYŻÓW COMPLEX						
				C	Czyżów Interstadial								
				a	Ferdynandovian Interglacial								
				b	Mazovian Interglacial - Holstenian								
				c	palaeosoil?								
KUCÓW FORMATION (glacial)					TILL 2B		ELSTERIAN						
					?								
					TILL 2			SANIAN					
					TILL 2A								
					TILL 2A								
					TILL 1		NIDANIAN						
					TILL 1								
ŁEKIŃSKO FORMATION					PRE - TIGLIAN		LOWER PLEISTOCENE						
PLIOCENE													

Petrographic features of the Saalian tills and neodymium isotope analyses were expected to help in the age interpretation of the tills found at the top of the Pleistocene deposits.

3 Materials and methods

3.1 Fieldwork in outcrop

The fieldwork mainly included lithological analyses, petrographic analyses, and macrofossil sampling of the deposits. They were exposed at mining levels I and II on the eastern wall (Figures 1–3). Glacial till samples were collected for lithopetrographic and geochemical analyses. Biological material

was helpful for setting boundaries of the Saalian complex. In 2014, seven till samples were taken (Nos. 465–471), and further five samples (Nos. 461–464a,b) were taken in 2016. A geological cross-section of the eastern wall at mining levels I and II, and a detailed sketch of exposed deposits were compiled in 2014 and 2016 (Figure 2).

3.2 Petrographic analysis

Petrographic analysis of gravels from the tills was performed on the 5–10 mm fraction, according to the standard specified by the Polish Geological Institute – National Research Institute in Warsaw [26]. A population of gravels (preferably more

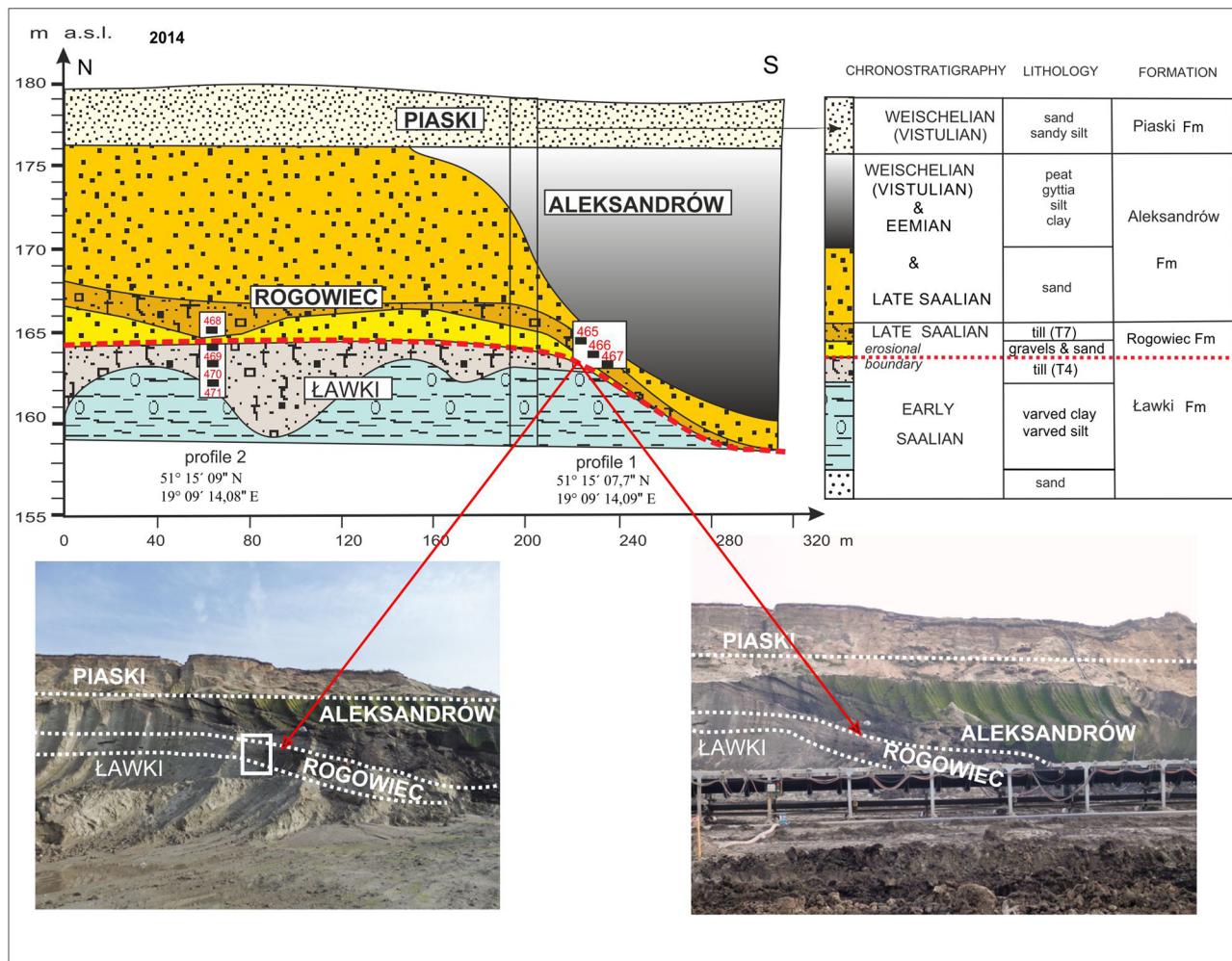


Figure 2: Geological cross-section of the Pleistocene sediments within the eastern part of the Szczerów outcrop (July 2014).

than 300 pieces) was divided into two groups. The first group included northern rocks from Scandinavia. The second one consisted of local rocks from the Polish Lowlands and those from near the Kleszczów Graben.

Petrographic coefficients O/K, K/W, and A/B have been determined according to the assumptions given by Gronkowska et al. [27]: O – total of sedimentary rocks [LPI + D + S + Sh]; K – total of crystalline rocks [Cr + Q]; W – total of carbonate rocks [LPI + D]; A – total of non-resistant rocks [LPI + D + Sh]; B – total of resistant rocks [Cr + Q + S].

The petrographic types of rocks are distinguished in the study, and their symbols are as follows:

- northern, Scandinavian rocks: Cr – crystalline (including schists); Q – quartz; LPI – Lower Paleozoic limestones; D – dolomites; S/Qt – sandstones and quartzites; Sh/ Ps – shales/Paleozoic slates;
- local rocks: L/M – limestones and marls; S – sandstones; M1 – Paleogene “older” mudstones; M2 – Neogene

“younger” mudstones; Q – quartz; Fl – flints; H – hornstones; F – phosphatic concretions; Pt – pyrite concretions; C – lignite/brown coal.

For full interpretation of petrographic coefficients of tills, the D/LPI ratio is also used, which relates the amount of dolomite clasts to limestone clasts in the carbonate group [3,28].

3.3 Neodymium isotope ratios

Rare earth element (metals) analysis was performed on sample Nos. 461–471, collected from deposits of the Rogowiec and Ławki Formations (see Table 4). The analyses were performed at Laboratoire Géodynamique et enregistrement Sédimentaire (PDG-REM-GM-LGS), Plouzané, France – using two methods [29,30]. ε_{Nd} values were used with the new Chondritic Uniform Reservoir (CHUR)

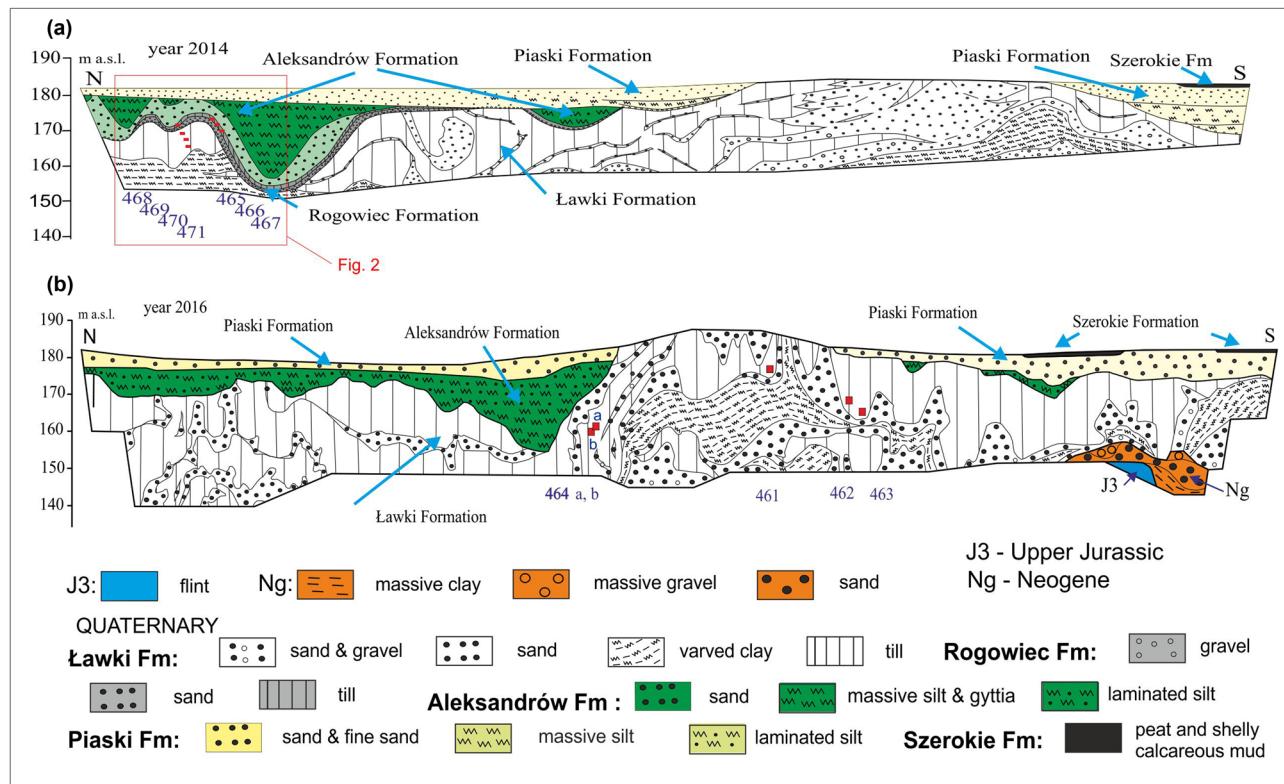


Figure 3: Geological cross-section of the quaternary sediments within the eastern part of the Szczerów outcrop (a: 2014 year, b: 2016 year).

value [29]. The ICP-MIS lab at IFREMER has started using this reference value. However, “ ε_{Nd} ” values are commonly presented using the “old” CHUR value of Jacobsen and Wasserburg [30]. Switching between these two reference values changes the resulting “ ε_{Nd} .” In the case of our data, this introduced a -0.15ε -value shift. Here, we report ε_{Nd} values using both CHUR values [29,30].

The ε_{Nd} values of our studied samples were published in Boswell et al. [31]. Replicate analyses of the JNd1 standard solution yield an estimated measurement uncertainty of $\pm 0.3 \varepsilon$ -units (2σ).

3.4 Macroremnant studies

A sample containing macrofossils (No. 464b) was collected from variously grained sands. The most frequent and common material, although difficult to examine, was organic detritus that has not been taken into consideration in microscopic analysis because the material was highly disintegrated.

First, the sample was sifted to find all, even minor, macrofossils. Then, if needed, macrofossils (or their minor fragments) were left in the mixture of glycerol and ethanol 100% (1:3) for a few days to soften the hard tissue [32,33].

To identify the plant genera, transverse, radial, and tangential hand-made sections were cut using a razor blade. To examine anatomical (diagnostic) details of macrofossils, the sections were examined in a BX50 Olympus microscope and documented using a DP71Olympus Camera cooperating with a Cell B Software (Olympus Optical Co., Poland). Greguss [34,35] and Schweingruber [36,37] keys as well as a Website: InsideWood 2004 [38] were used for the identification of wood genera. Terminology was based on the IAWA list of macroscopic features for hardwood and softwood identification [39,40]. The digital images were processed with Macromedia Fireworks MX2004.

4 Results

4.1 The main Saalian and Eemian horizons

The Middle Polish Complex is represented by deposits of the Ławki Formation (Odranian, Older Saalian) and the Rogowiec Formation (Wartanian, Late Saalian). The Ławki Formation deposits are variable lithologically. Glaciofluvial

Table 4: Neodymium isotope signature of the clay-silt fraction (<63 µm) of tills from the Szczerców outcrop

No.	Date of collection	Lithotype, formation	$^{143}\text{Nd}/^{144}\text{Nd}$ mean (norm)	$\pm 2\text{ se}$	ϵ_{Nd} with "new" CHUR value	ϵ_{Nd} with "conventional" CHUR value	$\pm 2\text{ se}$	$\pm 2\sigma$
465	2014	T7, Rogowiec Formation	0.511921	0.0000053	-13.8	-14.9	0.10	0.3
466	2014	T7, Rogowiec Formation	0.511900	0.0000053	-14.2	-14.9	0.10	0.3
467	2014	T7, Rogowiec Formation	0.511906	0.0000060	-14.1	-14.3	0.12	0.3
468	2014	T7, Rogowiec Formation	0.511913	0.0000068	-14.0	-14.0	0.13	0.3
469	2014	T4, Ławki Formation	0.511872	0.0000062	-14.8	-14.4	0.12	0.3
470	2014	T4, Ławki Formation	0.511875	0.0000038	-14.7	-14.3	0.07	0.3
471	2014	T4, Ławki Formation	0.511906	0.0000047	-14.1	-14.2	0.09	0.3
461	2016	T4, Ławki Formation	0.511958	0.0000050	-13.1	-13.3	0.10	0.3
462	2016	T4, Ławki Formation	0.511852	0.0000053	-15.2	-15.3	0.10	0.3
463	2016	T4, Ławki Formation	0.511961	0.0000038	-13.0	-13.2	0.07	0.3
464	2016	T4, Ławki Formation	0.511864	0.0000042	-14.9	-15.1	0.08	0.3
Mean:		0.511903		-14.2				

Source: According to Boswell et al. [31], slightly modified.

sands, tills (Figure 2), locally strongly deformed glaciectonically, and ice-dammed lake silty-sandy series have been identified [41,42]. The ice-dammed lake deposits were examined thoroughly in the "Parchliny C" site (see Figure 1) [43–45]. The sequence contained silts, clays, as well as varved clay, sometimes with extremely thick summer layers (up to 1 m), dropstones, massive or laminated diamictons, and fine-grained sands (Figure 2). Black massive clay contained freshwater molluscan and ostracod species. The Ławki Formation deposits are 20–40 m thick.

The top and base of the series was erosionally truncated (Figure 3). In the south of the outcrop, the Ławki Formation deposits (tills, sands, gravels, and muds) overlie Upper Jurassic limestones (J3) and Neogene sands. They are overlain by the Wartanian Rogowiec Formation deposits – tills and glaciofluvial sands and gravels (see Figure 2).

The Rogowiec Formation deposits are represented by tills (Figure 2) and fluvioglacial/glaciofluvial sands, 10–12 m thick (Figure 3) [41,44], which are found at different depths (4–20 m b.g.l.; 156–175 m a.s.l. [see Figure 2]). The contact with younger deposits is locally erosional (Figure 3) and partly sedimentary. The resulting depressions in the Wartanian till (kettle holes) are filled with lacustrine deposits of the Aleksandrów Formation from the Late Saalian, Eemian Interglacial, and Early Weichselian (Table 1). The Aleksandrów Formation deposits are represented by lacustrine muds, clays, and gyttjas, as well as peats (Figure 2). They have been found in the Leśna Niwa profile [46], near the villages of Grabek-Podlas and Ściegna-Stróża [41], as well as in the "Parchliny 2012" [47,48], "Parchliny 2014" (Figure 2) [23], and "Parchliny 2016" sections (Figure 3) [49]. These were benchmark intervals at the top of the Middle Polish Complex. The sedimentary sequence is terminated by Plenivistulian sands and muds of the Piaski Formation and Holocene sands and peats of the Szerokie Formation (Figure 3).

Profile "Parchliny 2016" (Late Wartanian to Early Eemian) is marked. Given the presence of sandy or muddy glacial rafts, the tills are suggested to be subdivided into units 1–4. This subdivision is only to indicate that the tills do not form a single stratigraphic horizon.

4.2 Petrographic analysis of tills

In 2014, till samples were collected near the "Parchliny 2014" section (see Figures 2 and 4 and Table 2). They came from two till horizons, separated by sands. It was found that they are characterized by high carbonate content (40–50%) as compared to other tills [24]. Both tills

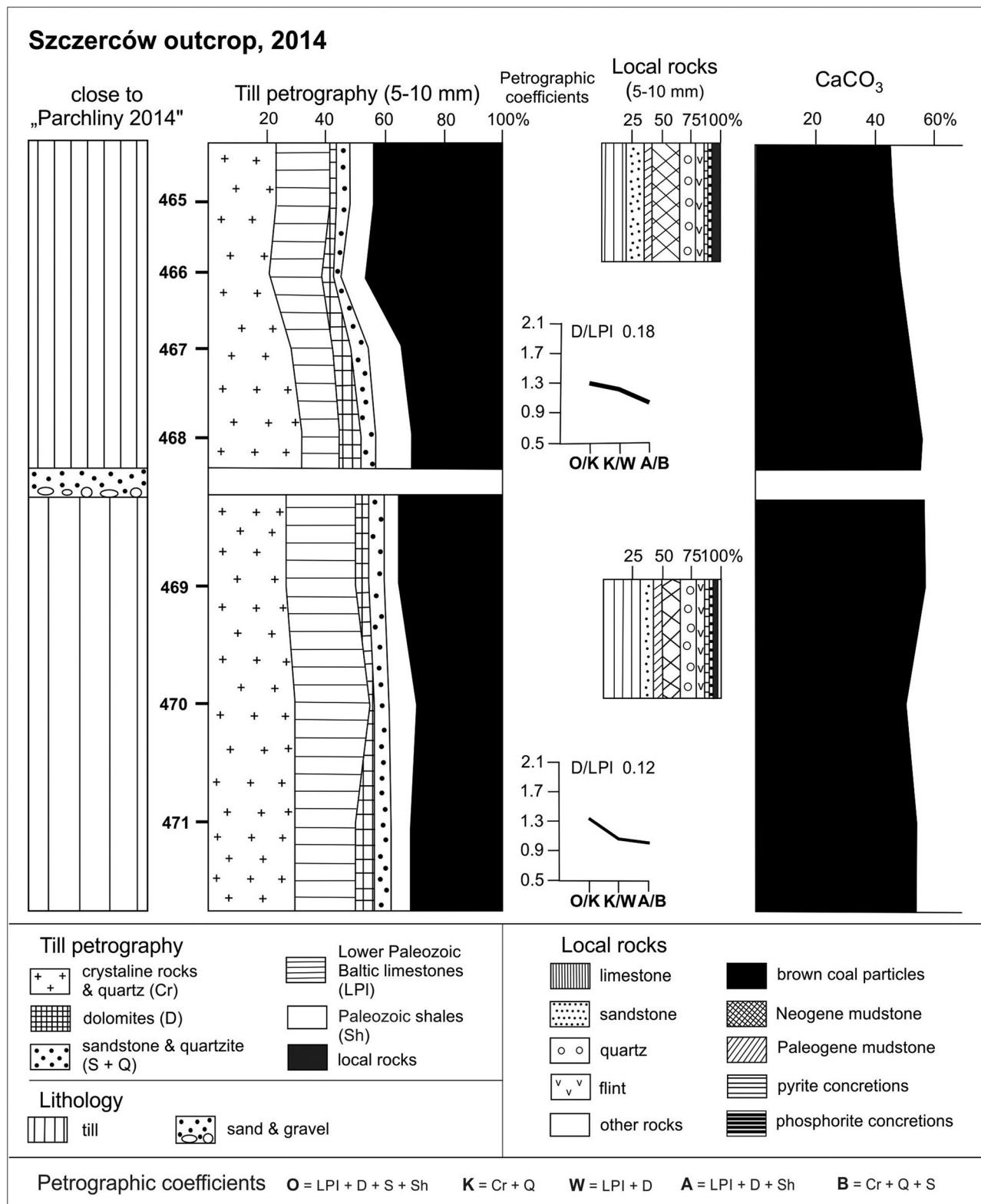


Figure 4: Results of petrographic analyses of till samples collected in 2014.

Table 2: Petrographic coefficients (O/K–K/W–A/B, D/LPI) of the upper tills from the Szczerców outcrop, obtained in 2014

Lithotype T7 (Rogowiec Formation)	Sample numbers:
1.29–1.23–1.04 (1.12–1.55)–(1.00–1.39)–(0.90–1.33)	465–468
D/LPI 0.18 (0.08–0.28)	
Lithotype T4 (Ławki Formation)	Sample numbers:
1.33–1.06–1.01 (1.29–1.36)–(1.04–1.10)–(0.99–1.03)	469–471
D/LPI 0.12 (0.02–0.21)	

Note: A range of variation is given in brackets, in bold – mean values.

contain a considerable amount of local (24–44%) and Scandinavian (22–35%) rocks. The group of local rocks in the upper till is dominated by equal amounts of limestones and Neogene mudstones (ca. 20%), whereas the lower till displays a different proportion of these rocks (limestone 30% and Neogene mudstone 15%). The contents of Baltic limestones and dolomites differ significantly. In the Odrianian tills, the content of Baltic limestones is higher than in the Wartanian tills. The amount of dolomites, in turn, is greater in the upper tills of the Rogowiec Formation than in the Ławki Formation tills.

At the current stage of geological investigations, the upper till is tentatively correlated with lithotype T7 (Rogowiec Formation), and the lower till with lithotype T4 (Ławki Formation). This lithostratigraphic correlation is performed bearing in mind that the results of petrographic coefficients of the tills do not match those obtained a couple of years ago [24]. In a sense, this correlation is determined by the geological setting of deposits, especially those at the “Parchliny 2014” section, which are included in the Aleksandrów Formation (Figure 2).

The next till samples were acquired for analyses in 2016 (see Figure 3 and Table 3). One of them (No. 464) came from a palaeolake – from the “Parchliny 2016” site – and further three samples (Nos. 461, 462, and 463) were collected more to the south, within a zone deformed

glaciectonically and tectonically. Sample Nos. 461 and 464 correspond to lithotype T4 (Ławki Formation). Given the location of sample No. 464 underneath the palaeolake (filled with the MIS6/MIS5e boundary deposits in its lower part), the Rogowiec Formation lithotypes were expected to be found. In turn, the coefficients of sample Nos. 462 and 463 are different, although the samples are taken from sites located close together, and they are also different from those recorded for sample Nos. 461 and 464. These tills were deformed, and the magnitude of the deformation is manifested by lithological variability, steep strata dips, and local high thicknesses (see Figure 3).

4.3 Neodymium isotope ratios

Currently, rare earth element (metals) analysis has been included in the range of till investigations. It covers the tills from sample Nos. 465–471 and 461–464 collected from deposits of the Rogowiec and Ławki Formations. For these samples, the average $^{143}\text{Nd}/^{144}\text{Nd}$ value is 0.511903 (0.511852–0.511961) and the average ε_{Nd} is –14.2 (–15.2 to –13.0) (see Table 4). The mean ε_{Nd} signatures for the Saalian sediments are -14.4 ± 0.7 ($n = 11$), respectively.

The difference in ε_{Nd} between the sample Nos. 461–464 does not appear to have any discernible relationship with the 5–10 mm till petrography, as inferred from Figures 4 and 5 attachment. Given the similarity in grain size distribution of the studied samples, with pronounced modes in the coarse silt fraction, we expect that the ε_{Nd} signature of the tills is controlled by the provenance of heavy minerals (e.g., monazite and allanite) and not by other minerals (e.g., feldspars, pyroxenes, quartz, and clays) [50,53].

In regard to the Rogowiec samples, note that these samples cluster much more closely around the measured ε_{Nd} value for the Vistula River, with sample No. 467 being somewhat of an exception (more non-radiogenic).

Table 3: Petrographic coefficients (O/K–K/W–A/B, D/LPI) of tills from the Szczerców outcrop, obtained in 2016

Lithotype T4 (Ławki Formation)	Sample numbers: 461 and 464
0.89–1.27–0.78 (0.80–0.97)–(1.12–1.41)–(0.70–0.87)	
D/LPI 0.04 (0.02–0.06)	
1.08–1.03–0.93 D/LPI 0.09; 0.38–3.20–0.32 D/LPI 0.25; Mean 0.73–2.11–0.62 D/L 0.17 Mean 0.81–1.69–0.70 D/LPI 0.11	Sample numbers: 462 and 463

Note: A range of variation is given in brackets, in bold – mean values.

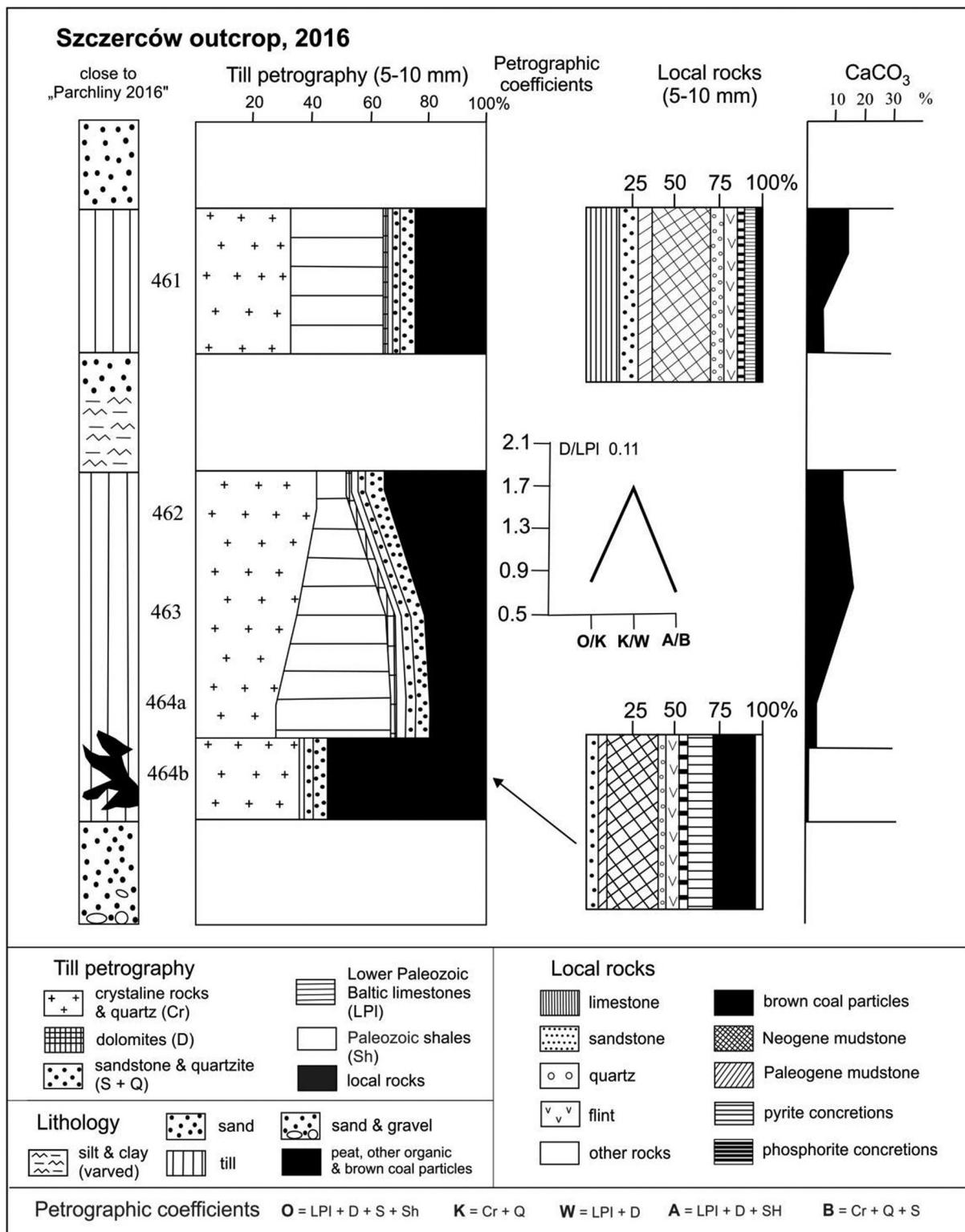


Figure 5: Results of analyses of till samples collected in 2016.

4.4 Results of macroremnant studies

The main aim of the analyses of the macrofossils was the identification of the plant genera to which the wood fragments belong. Therefore, 18 woody remnants were examined.

Five of them were identified as gymnosperm wood characterized by the presence of distinct resin canals with thin-walled epithelial cells. On the radial section, the heterogeneous rays were visible, consisting of parenchyma cells and ray tracheids located on the ray margins. The characteristic large fenestriform pits, on the contact between ray parenchyma cells and axial tracheids, were also visible. These features indicate the wood as pine (*Pinus*) wood (Figure 6a).

The next five fragments were marked as willow wood (*Salix* sp.) based on the presence of the following features: (1) diffuse-porous wood with a radially arranged group of two to five vessels (Figure 6b), (2) simple perforation plate of vessel elements, (3) uniserrate rays, and (4) heterogeneous rays. The last feature was recognized in a radial section, where the marginal cells of the rays with large simple pits with four to eight pits in height occurred (Figure 6c).

Three other woody macrofossils were identified as a birch (*Betula* sp.) wood. This genus is characterized by diffuse-porous wood with a radially arranged group of three to five vessels (or more); the scalariform perforation plates in the vessels elements (Figure 6d) and bi-tetraserrate rays (Figure 6e). Five macrofossils could not be identified. They were fragments of young stems of dicotyledonous trees, 1–2 years old, and heavily compressed.

5 Discussion

The results of the research conducted in the Szczerów outcrop supplement the data on the Pleistocene section in central Poland (Table 1). Although no continuous horizons of organic sediments have been found in the Czyżów Formation (Czyżów Complex, Holsteinian/Ferdynandovian [see Table 1]), the studies of macroremains from the bottom of the Middle Polish Complex helped to determine the lower limit of the Saalian complex (see also [54,55]). The investigated sandy deposits contained macroremains that originated in the valley bottom environment. The deposits are

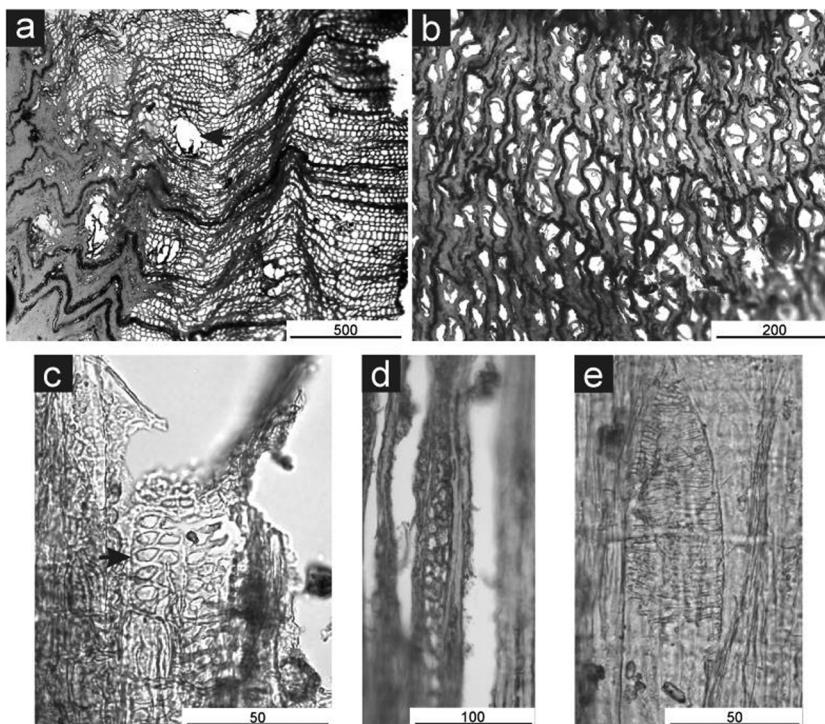


Figure 6: The anatomical features of the woody macrofossils. (a) Transverse section of *Pinus* wood; the resin canals with thin-walled epithelial cells occur (arrows). Hardly compressed wood is visible on the left side of the image. (b) *Salix* sp. wood; (c) marginal cell of the ray with large simple pits (arrow) between the ray and vessel shown in radial section; usually four to eight pits in height occur in the upright, marginal cell of the heterogeneous ray; (d and e) *Betula* sp. wood; (d) tangential section showing a triseriate ray; (e) scalariform perforation plates are presented in radial section. Scale bar in μm .

compressed probably as a result of the pressure of overburden strata and compaction. They form a discontinuous horizon and overlie both the Neogene deposits and probably the partially preserved sandy deposits of the South Polish Glaciations. Their stratigraphic position and preserved remains of deciduous and coniferous trees are ambiguous. The deposits might have been accumulated in both the Pliocene and the interglacials or interstadials of the Lower or Middle Pleistocene (Czyżów or Łękińsko formations). They were extensively described in the nearby Bełchatów [1]. The sandy deposits with macroremains are covered by relatively thick sandy and loamy sediments (Figure 3).

The Middle Polish Complex (Saalian) starts with the Ławki Formation deposits. In the study area, muddy-clay series and a single till horizon (lithotype T4) were found in 2014–2016. The till is strongly deformed and rests on Jurassic, Neogene, and Early Pleistocene deposits (Figure 3). In the Bełchatów outcrop, like in the Szczerców outcrop, these deposits were strongly deformed. In the nearby Bełchatów mine field, they were found at the top of the lower structural level, disturbed as a result of tectonic processes [1]. The deposits investigated in the Szczerców outcrop may have been deformed because of tectonic Chabielice fault (which uplifted Jurassic limestone), glaciotectonic (on time of the advancing Wartanian ice sheet), and halotectonic factors (i.e., DSD).

The ε_{Nd} values of the studied Saalian tills are also compared to 15 samples obtained from the Greater Poland region, collected from glacial or glaciofluvial deposits accumulated during the last glacial maximum (LGM) [51]. Compare the ε_{Nd} of these Szczerców tills with: (1) sediments from the mouth of the Vistula River ($-14.3 \pm 0.3, 2\sigma$) and (2) Weichselian glaciogenic sediments from Poland ($-15.1 \pm 0.8, n = 15, 2\sigma$). Samples 461 and 463 are more radiogenic (younger in the Sm/Nd sense) than any of the 15 samples.

For these samples, the average $^{143}\text{Nd}/^{144}\text{Nd}$ value is 0.511855 (0.511792–0.511921), and the average for ε_{Nd} is -15.2 (-16.5 to -13.9). The Rogowiec samples are not glaciotectonically disturbed. These samples display a similar relationship between ε_{Nd} and grain size as the Weichselian samples used by Boswell et al. [52] to infer sediment transport distances by the Baltic Ice Stream. It is plausible that this relationship could hold for the Saalian Baltic Ice Stream(s) as well as the findings of Boswell et al. [31].

The ε_{Nd} values of the studied Saalian tills are also compared to samples from the Greater Poland region, collected from glacial or glaciofluvial deposits accumulated during the LGM ($-15.1 \pm 0.8, n = 15, 2\sigma$) and the mouth of the Vistula River ($-14.3 \pm 0.3, 2\sigma$) [51]. Analysis of ε_{Nd} values within the studied sections reinforces our lithostratigraphic

correlations and reflects the differences in disturbance between the Ławki and Rogowiec Formation tills.

Samples 461 and 463 are more radiogenic (younger in the Sm/Nd sense) than any of the 15 Polish samples analyzed by Toucanne et al. [51]. This 2 ε -unit difference between samples 461/463 and 462/464 evinces a lack of homogenization of the Ławki Formation till at this site. Because the samples are disturbed, it is possible that samples 461 and 463 have incorporated relatively radiogenic fine sediments (silts and clays) from an unidentified, presumably local source inconsistent with the expected signature of glaciogenic sediments from the Weichselian glaciation (LGM) ($-15.1 \pm 0.8, n = 15, 2\sigma$) or the Vistula river background ($-14.3 \pm 0.3, 2\sigma$).

The Rogowiec tills are not disturbed in the Szczerców outcrop, and this is reflected in the dispersion of these samples' ε_{Nd} values. The standard deviation in the ε_{Nd} of the Rogowiec samples (both Early and Late Saalian, but also separately) is much less than for the Szczerców (i.e., ± 0.36 vs ± 1.15). The increased dispersion in ε_{Nd} of the Ławki Formation tills reflects their disturbed (tectonically, glaciotectonically, or halotectonically) nature and is likely explained by the incorporation of radiogenic fine sediments of local provenance.

The undisturbed Rogowiec samples display a similar relationship between ε_{Nd} and grain size as the Weichselian samples used by Boswell et al. [52] to infer sediment transport distances by the Baltic Ice Stream. It is plausible that this relationship could hold for the Saalian Baltic Ice Stream(s) as well [31].

The Ławki Formation deposits are erosionaly overlain by the Rogowiec Formation deposits of the Wartanian Stadial. Three till horizons (lithotypes T7, T6, and T5) have been identified in the Bełchatów outcrop and included in the Rogowiec Formation. One lithotype of till from this formation, most likely T7, has been recognized so far in the Szczerców outcrop. Initial soil horizons have also been identified within the Rogowiec Formation in the Bełchatów outcrop. This would suggest longer breaks in glacial deposition and even decay of the ice sheet over a larger area, so that the conditions allowing soil-forming processes could take place. In the Bełchatów outcrop, these tills are not deformed. Wherever there are any disturbances, they are visible only in the basal till (T5) and are represented by deformation resulting from the stress of the advancing Wartanian ice sheet [8]. In the case of the till filling the Kleszczów Graben in the western sector, there are various types of deformation (see Figure 3) of different origins, like in the case of the Ławki Formation deposits. In the hinterland and foreland of both outcrops, there is a single till horizon of the Wartanian Stadial/Glacial. In

depressions, the Wartanian sands and tills overlie lacustrine gyttjas and peats of the Eemian Interglacial.

In the “Parchliny 2012” [47], “Parchliny 2014” [48], and “Parchliny 2016” [49] sections, Eemian Interglacial deposits, deposited as a result of lake basins infilling with terrigenous sediments, have been identified. Organic deposits were of key importance for determining the lithostratigraphic position of the underlying tills. The “Parchliny 2014” and “Parchliny 2016” sections also encompass the MIS6/MIS5e and MIS5e/MIS5dcba transitions. These paleolakes developed at the end of MIS 6, later they were probably connected by subglacial channels, and even later partly used by flowing waters and transformed by them [22]. This was the case with the Krasówka stream, whose “valley” is too large compared to its catchment area. Some of these studies are still underway, but it may turn out that Lower Vistulian deposits, previously unknown in this area (“Parchliny 2014” site), will be found.

6 Conclusions and recommendations

1. Organic horizons defining the boundaries of Saalian deposits have been found – at the base, macroremains of trees in sediments of unclear stratigraphic position (Pliocene, the Lower Pleistocene Łękińsko Formation or the Czyżów Formation (with Czyżów Interstadial; Holsteinian). At the top, there are gyttjas and peats of the Aleksandrów Formation (Late Saalian, Eemian, Early Weichselian), which accumulated in depressions of the Wartanian surface.
2. The Saalian deposits are represented by clay-muddy and sand-gravelly sediments and two till beds T4 and T7 corresponding to the Odrianian (Ławki Formation) and the Wartanian (Rogowiec Formation). These tills are strongly disturbed. The possibility of deformation as a result of tectonics within the Kleszczów Graben (Chabielice fault), salt tectonics of the DSD, and glaciotectonics is pointed out.
3. Lithopetrographic studies and rare earth element (metals) analyses helped distinguish the features of these tills, wherever deformed, and compare them with similar ones examined in central Poland. The petrographic coefficients for lithotype T4 (Ławki Formation) are 1.33–1.06–1.01 and 0.12 (2014) and 0.81–1.69–0.70 and 0.11 (2016), and for lithotype T7 (Rogowiec

Formation) these are 1.29–1.23–1.04 and 0.12 (2014) and 0.89–1.27–0.78 and 0.04 (2016). Investigations of the Nd isotopic compositions of the studied samples allow reinterpreting of till deformation and tentative lithostratigraphic correlations. Neodymium isotope ratios prove that ε_{Nd} values less than –12.4 reflect the southern Fennoscandian Ice Sheet sediment provenance, but ε_{Nd} values greater than –12.4 indicate western European origin. “P” (-14.4 ± 0.7) and “D” (-12.4 ± 0.3) ε_{Nd} values (2σ) correspond to mean signatures of Saalian glaciogenic moraine sediment from Poland and Denmark, respectively.

4. The authors believe that in the future more emphasis should be placed on the study of tills. As the mining operations are in progress, this work has to be coordinated with the analysis of geological and mining sketches of walls of the Szczerów outcrop. As the mining works are getting closer to the DSD, the deformation of Pleistocene formations is expected to be greater. Only the combined data resulting from both lithopetrographic studies of the tills and analysis of their distribution in the vertical section (cross-sections analyzed in time) can bring us a better picture of the changes that occurred in the past.
5. Perhaps the issue of stratigraphic subdivision of the Saalian tills in both outcrops should be considered another way. When studying the Pleistocene formations, a relatively complete section is assumed to occur here, because the Kleszczów Graben is treated as a kind of trap that was to preserve this section. However, if this assumption were put aside, it would seem that the two major glacials, which have left behind the four till horizons, must have been separated by at least one interglacial, although no interglacial deposits have been preserved. This could have been the result of processes of glacial exaration, glaciotectonics, and erosion of glacial and river waters.

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